

Our aim is to test a range of geodynamic models that can explain the uplift of the Colorado Plateau, if uplift there was. Many mechanisms have been proposed and summarized in the literature [e.g. McGetchin 1980]. The main mechanisms fall in three different categories: (a) Late Cretaceous to early tertiary uplift related to Sevier-Laramide orogeny (80 to 40 Ma) [e.g. McQuarrie and Chase 2000]; (b) mid-Tertiary uplift related to removal of flat subduction (either through mechanical thinning of continental mantle lithosphere and subsequent removal of flat slab [Spencer 1996], release of negative dynamic topography [Mitrovica et al 1989, Gurnis 1992] or hydration of the lithosphere from volatiles derived from the Farralon slab [Humphries et al 2003]) and (c) Late-Tertiary uplift associated with regional extensional tectonism, either by large removal of instable lithosphere [Bird 1979] or heating from below. We propose to test these geodynamic models using the codes available through the Computational Infrastructure for Geodynamics framework.

We use a mantle convection model, so-called CitcomT CitcomT is a solving the equations of momentum, cor transport equations :

> $u_{i,i} = 0,$ $-P_{,i} + (\eta u_{i,j} + \eta u_{j,i})_{,j} + Ra\delta\rho\delta_{ir} = 0,$ $(\delta\rho)_{,t} + u_i(\delta\rho)_{,i} = (\delta\rho)_{,ii},$

in 3D and in spherical coordinates.

It computes dynamic topographies using the boundar [Zhong et al 1993].

The viscosity varies laterally and is temperature-pressu and radial viscosity variations can be included in the n the effective viscosity to the desire value in a given rec

The code to account for regions that are chemically di to simulate continental lithosphere/tectosphere. It is this implementation includes the computation of the a chemically different layer.

We present three models which test the hypothesis of Colorado Plateau due to removal of the the flat subdu we include a buoyant continental tectosphere that is a subduction. In each model, decoupling between the s and the overiding lithosphere is ensured by including viscosity between the two. Model 1 simulates an abno lithosphere below the plateau, Model 2 assumes a thir Model3 is equivalent to model 2 except that the zone of larger.

Geodynamic models for the "uplift" and erosion of the Colorado Plateau

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Γ[e.g. Billen et al 2003] ntinuity and	Model 1 Temperature
	Temp (C)
	Temp (C) $\int_{-200}^{-200} \int_{-200}^{-400} \int_$
	Femp (<i>b</i>)
ary flux method	Model 2 Temperature Temp (C)
	Lime 3.1125 Myrs -200 -400 -600 -600 -600 -1000 -1000 -12
ure dependant. Lateral model, by setting gion.	-1500 -1000 -500 0 500 1000 1500 Temp (C) 1000 -200 -400 -400 -600 -1000 -500 0 500 1000 1500 Temp (C) 1000 10
ifferent [Moresi 1998] important to note that advection of the	time 19.7125 Myrs Temp (C) -200 -400 -400 -600 -600 -600 -1000 -500 0 500 1000 1500
	Model 3 Temperature
f uplift of the action. In each model, shaved by the flat subducting plate g a region of low ormally thick oceanic nner oceanic plate and of decoupling is made	time 2.5938 Myrs Temp (G) -200 1200 -400 1000 -600 - -800 - -100 - -1500 -1000
	time 8.8188 Myrs Temp (©) -200 -200 -400 -200 -600 -200 -600 -200 -600 -200 -1000 -200 -1000 -500 0 500 1000 1500
	time 19.7125 Myrs -200 -400 -400 -600 -400 -1000 -1500 -1000 -500 0 500 1000 1500 -1000 -500 - -50







which simulates the evolution of hillslope and river incision. We present an example of a simulation.

The model starts with a base level drop. Consequently, erosion is induced at the edges of the plateau. This, in turn, causes isostatic uplift on the edges of the plateau, slowing down erosion along these scarps and forming an internally drained basin. Ultimately, the basin will overfill and lead to rapid erosion along the Grand Canyon.

Computations were perfomed on the Pangu facilities at the Geological and Planetary Sciences Division, California Institute of Technology