

Figure 2: The North Anatolian fault and the western migration of events in 20th century. The Mw7.4 Izmit earthquake occured on August 17, 1999 (1999a) and Mw7.4 Duzce earthquake occured on Nov 12, 1999 (1999b). Both earhtquakes created significant surface offset and

Mw 7.1 Duzce Earthquake: Multiple datasets and implications about rupture kinematics A. Ozgun Konca¹, Sebastien Leprince¹, Jean-Philippe Avouac¹, Don Helmberger¹ ¹Tectonic Observatory, California Institute of Technology e-mail: ozgun@gps.caltech.edu

4. Modeling of strong-motion, GPS and InSAR vr=2 vr=3 vr=2.5-4 0 -10 0 10 20 -20 -10 0 10 20 -20 -10 0 10 15 20 25 direction. We have searched for the best fit rupture velocity, to confirm whether supeshear rupture velocities occured to the east. The slip models with different fixed rupture velocities and varying rupture velocities show a very consistent slip pattern due to constraints coming from geodesy. vr=3 -20 -10 0 10 20 -20 -10 0 10 20 8: The slip map (top) and rise 5. Fits to the strong-motion data GOL DZC mm. - Mm BOL 3 km/s 🔨 10 20 0 10 20 0 10 2 E N Z M-MM DZC away from the hypocenter. -Man Mon Man Mon Mon



obtained from SPOT imagery. observations.



6.Map view of slip and snapshots of rupture

Supershear rupture is only significant in east

The rupture starts subshear, accelerates to supershear speeds and slows down eventually.



Figure 10: Map view of total slip with time contours and snapshots of for slip every 2 second time windows



7. Fits to the GPS and InSAR data



8. Reconciling teleseismic and Strong

motions finding of this study is that in order to predict the teleseismic waves from joint inversion of strong-motions and geodetic data, we have to shift teleseismic waves by ~2 seconds. The shifts imply that when earthquake starts small, then the teleseismic picks can be later than the actual arrival time, leading to bulging around the hypocenter, while usually more slip happens

Figure 12: Telesesismic data (black) and forward prediction of the data from the joint strong-motion and geodetic inversion shown in Figure 10 (red). Greens are the fits from a teleseismic and geodetic joint inversion

P SUM 194	21.6	WMQ 66	244.7		strong motion and geodesy
P URI <u>166</u>	12.6	SH MAJO 77	107.3	_	teleseismic and geodesy
P SEY 147 50	25.3	SH YSS 41 73	102.8	SH ALE 350 50	287.6
PSI 102	25.1	SH YAK 33 58	94.0	SH FFC <u>334</u> 76	151.0
	39.5	SH MA2 28 68	129.4	SH SSPA <u>312</u> 76	100.1
P SA 83 49	53.5	квs <u>354</u>	9.6	SH SJG <u>288</u> 83	101.4
P MQ <u>66</u>	65.0	FFC 334	12.1	SH TSUM <u>194</u> 61	175.3
Puo 51	28.5	SSPA 312	18.8	SH LBTB <u>185</u> 65	197.1
$\frac{41}{73}$	31.3	P SJG 288 83	18.4	SH FURI <u>166</u> 32	286.1
Р АК <u>33</u> 58	34.4	P RCBR 250 76	25.3	CHTO 61	243.4
A2 28	28.2	ASCN 231	37.7	SH LSA <u>83</u> 49	209.1
LL $\frac{16}{66}$	26.3	Р ВGCA <u>201</u> 37	34.4	TATO 70 74	186.6
-10 0 10	20 3	0 -10 0	10 20 30) -10 0	10 20 3

Summary and Conclusions

1. We have modeled the Mw7.1 Duzce earthquake using an accurate 4 segment fault model

2. The Duzce rupture has extended further east than the models based on field

3. We have modeled the earthquake with all available data.

4. In Duzce Earthquake, rapid variations in rupture velocity are required to explain the strong-motion data, along with geodetic data

5. Supershear rupture velocity is local and only to the east of the rupture.

6. The teleseismic data can miss the beginning of the events leading to more compact models with a lot of slip around the hypocenter than the actual sources.