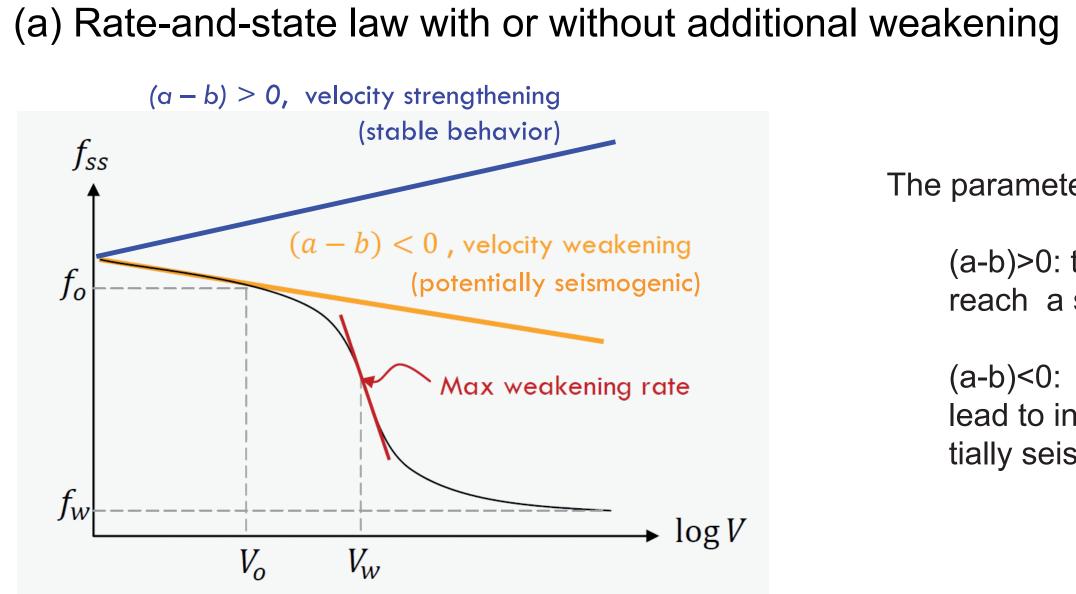


### Abstract

A number of factors which might be inferred from geological and geodetic observations are thought to influence fault seismic behavior: they include lithology which might control mechanical properties of fault zone, pore pressure and faults geometry. In theory the imfluence of these factors might be estimated from theoretical fault models and computer simulations. This is computationally challenging because this kind of study requires a proper account of the effect of these factors on rupture dynamics, at the scale of individual seismic event, as well as long sequence of rupture to capture the stochastic behavior of faults systems. In such studies it is computationally advantageous to not incorporate full inertial effects during simulated fast slip. That is why so-called quasi-dynamic methods have become increasingly popular, which approximately account for inertial effects (and hence seismic radiation) during simulated earthquakes through a radiation damping term. Such methods allow continuing simulations through the seismic phase, without having to pay significant additional memory and computational costs associated with modeling true wave-mediated effects. However, the resulting seismic events tend to have much slower slip velocity and rupture speeds and may modify significantly the resulting seismic events and hence the long-term fault behavior.

In this study, we compare the results of quasi-dynamic and fully dynamic simulations, with wave effects during simulated earthquakes. We consider the long-term fault behavior in two problems: (i) interaction of two velocity-weakening regions separated by a small velocity-strengthening patch and (ii) segments with additional pronounced rate-weakening during seismic slip. We find that, in the absence of additional seismic weakening, the two methods generally result in the same qualitative behavior, with similar slip patterns, although there are quantitative differences. However, in simulations with additional rate weakening, the two methods produce qualitatively different long-term results, with different slip patterns and significantly different levels of shear stress on the fault. Our eventual goal is to determine the range of applicability for the quasi-dynamic approaches.

### Numerical model statements



The parameters "a" and "b" in this equation allow describing 2 kind of material:

(a-b)>0: the friction increase with velocity, so no instabilities occurs and we reach a stable sliding. It is called the velocity strengthening behavior

(a-b)<0: the friction decrease with velocity, so an acceleration occurs, that lead to instability. It is called the velocity weakening behavior and it is potentially seismogenic

In those models we are using laboratory derived friction laws.

$$\tau = \bar{\sigma} \left( f_0 + a \ln \frac{V}{V_0} + b \ln \frac{V_0 \theta}{L} \right) = \bar{\sigma} f_{R\&S} \qquad \dot{\theta} = 1 - 1$$

R&S law is valid at low velocity. At seismic velocity, some additional weakening mechanism, au = like flash heating, can help the rupture to propagate.

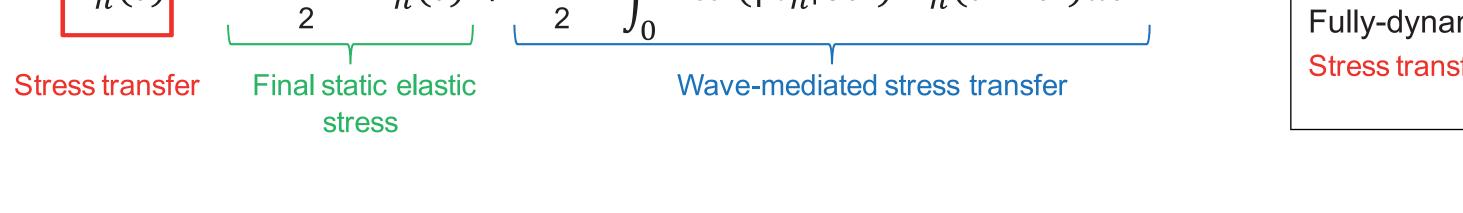
### (b) Quasi-dynamic versus Fully-dynamic modeling

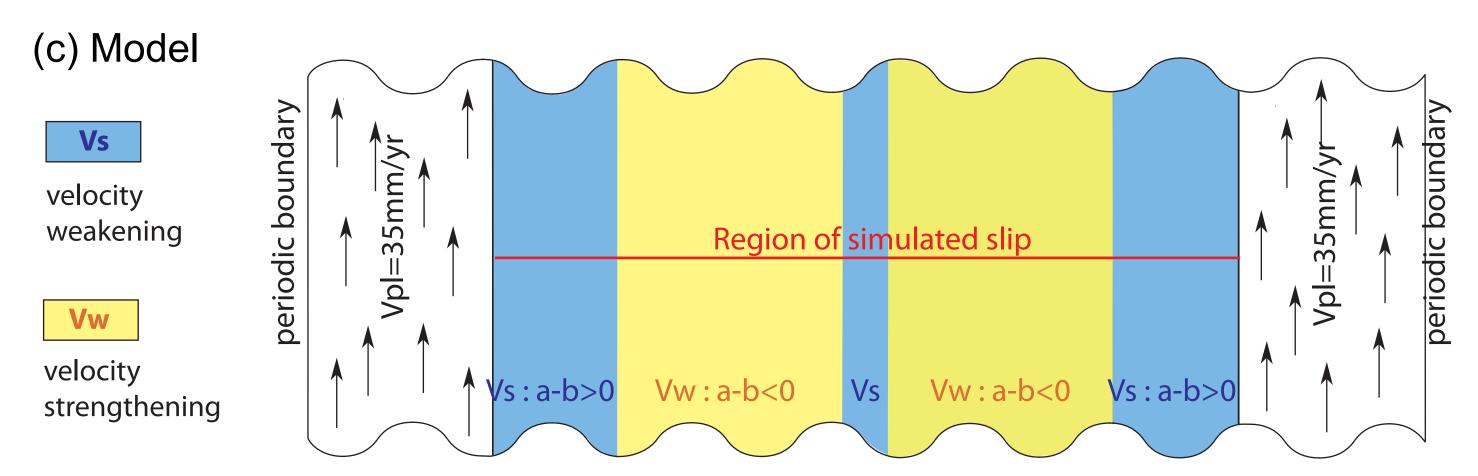
To solve the equations, model uses a Spectral Boundary Integral Method

$$\frac{\tau(x,t)}{\text{Shear stress loading}} = \frac{\tau^o(x,t)}{\sum_{\text{Shear stress loading}}} + \frac{f(x,t)}{2c} - \frac{\mu}{2c} V(x,t)$$

$$f(x,t) = \sum_{n=-N_{ele}/2}^{N_{ele}/2} F_n(t) e^{ik_n x} \text{ with } k_n = \frac{2\pi r}{\lambda}$$
Stress transfer in Fourier Domain
Evolution of stress during the rupture:
$$\int_{1}^{\infty} \int_{0}^{\text{Slip}} V(|k_n|ct') \dot{D}_n(t-t') dt'$$

$$F_n(t) = -\frac{\mu |k_n|}{2} D_n(t) + \frac{\mu |k_n|}{2} \int_{0}^{T_w} W(|k_n|ct') \dot{D}_n(t-t') dt'$$
Fully-dynamic formulation:
$$F_n(t) = -\frac{\mu |k_n|}{2} D_n(t) + \frac{\mu |k_n|}{2} \int_{0}^{T_w} W(|k_n|ct') \dot{D}_n(t-t') dt'$$





# Fully-dynamic vs Quasi-dynamic simulations of slip accumulation on faults with heterogeneous friction properties

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Allows simulating earthquake cycles their entirety, from accelerating slip in slowly expanding nucleation zones

to dynamic rupture propagation to post-seismic slip and interseismic creep

to fault restrengthening between seismic events.

$$= \overline{\sigma} \left( \frac{f_{R\&S} - f_{W}}{1 + \frac{L}{V_{W}\theta}} + f_{W} \right)$$

## Stress transfer = Final static elastic stress +

Wave-mediated stress transfer

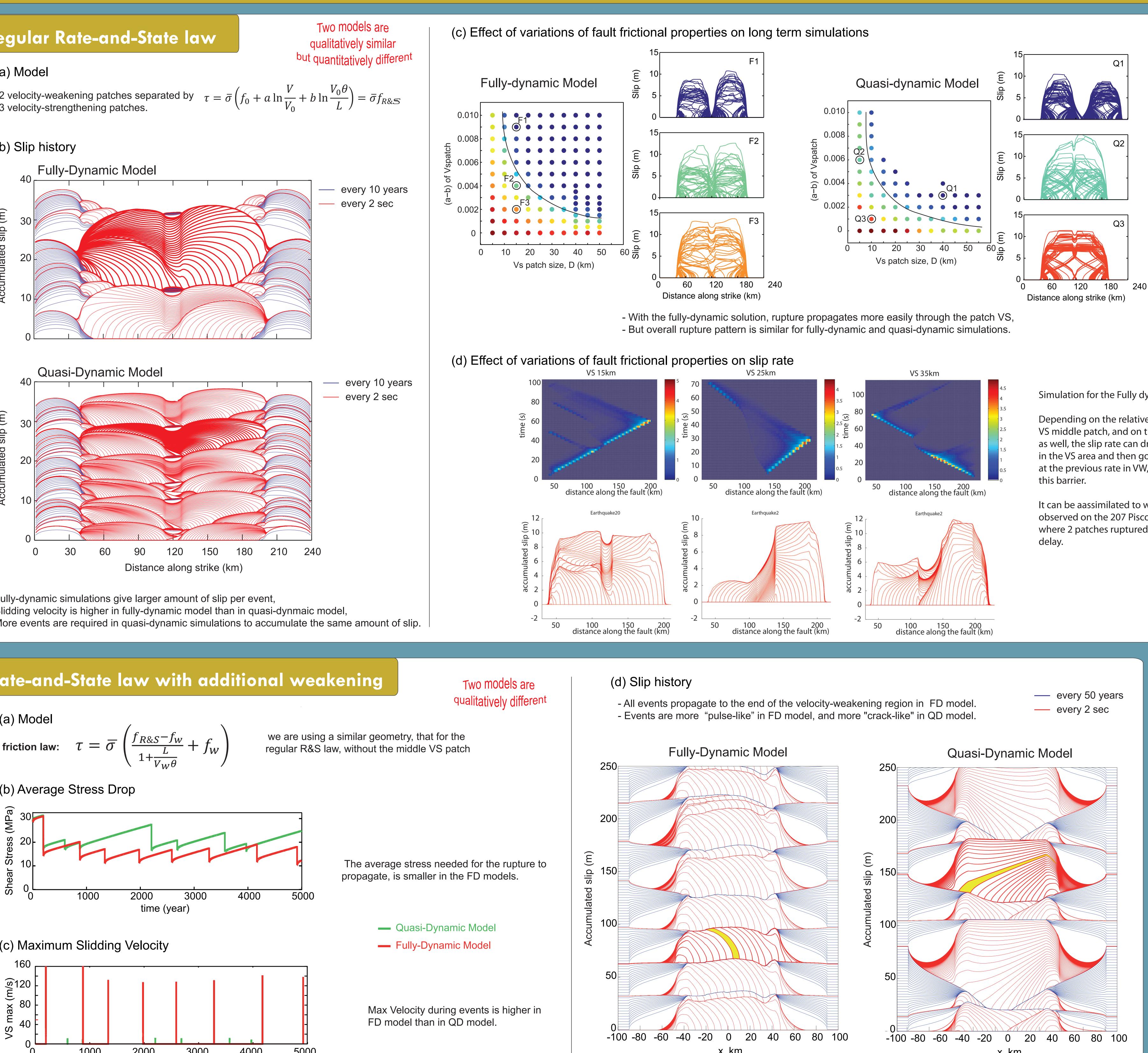
Model:

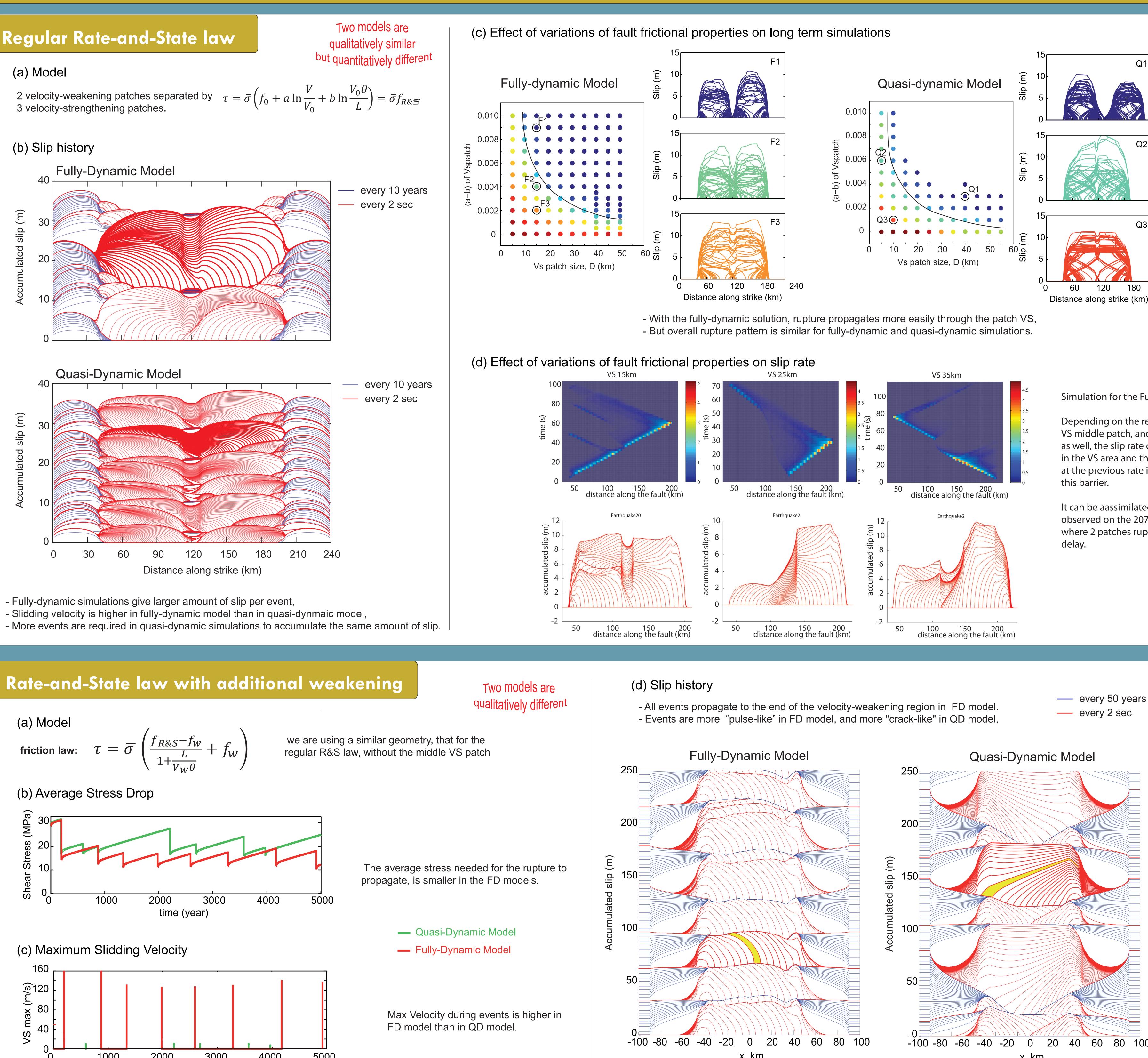
- 2D antiplane model with 1D fault - Equations solved for an infinite,
- uniform, isotropic, elastic space, - 2 velocity-weakening patches separated by 3 velocitystrengthening patches.

### Regular Rate-and-State law

### (a) Model

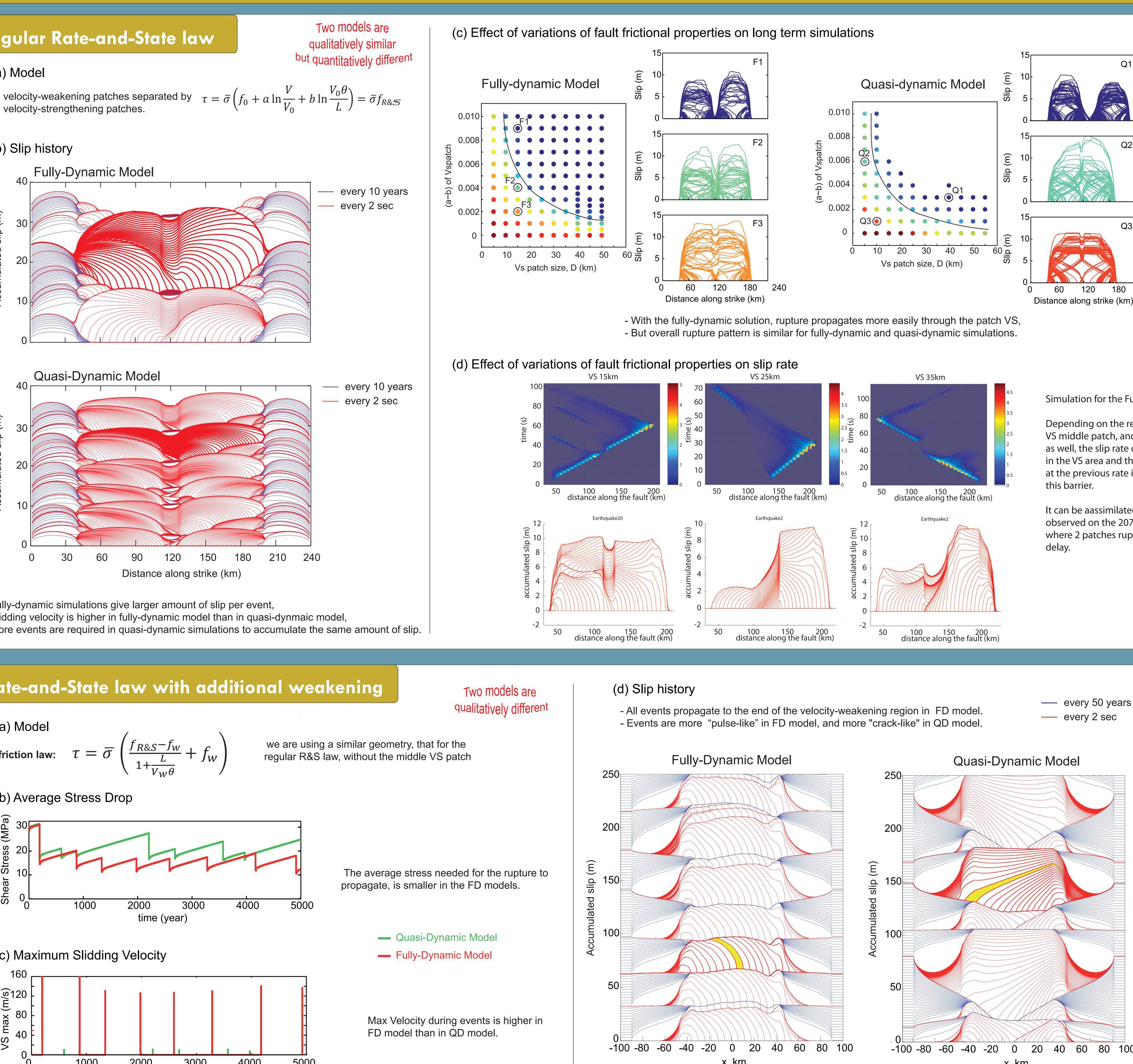
3 velocity-strengthening patches.

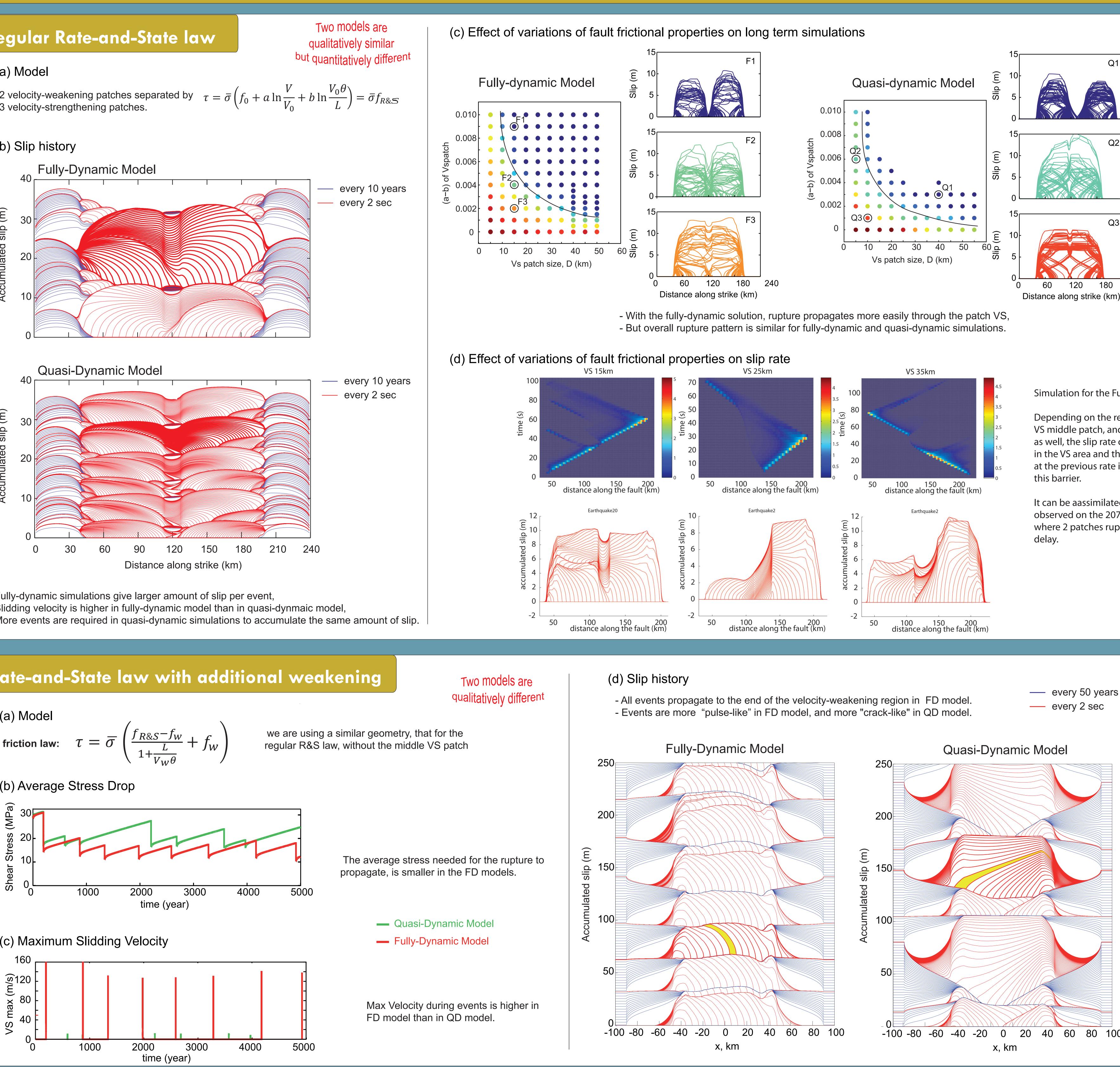




(a) Model

$$= \overline{\sigma} \left( \frac{f_{R\&S} - f_{W}}{1 + \frac{L}{V_{W}\theta}} + f_{W} \right)$$







Simulation for the Fully dynamic case.

Depending on the relative size of the VS middle patch, and on the pre-stress as well, the slip rate can drop drastically in the VS area and then goes up again, at the previous rate in VW, after crossing

It can be aassimilated to what has been observed on the 207 Pisco earthquake, where 2 patches ruptured with a 40 sec

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