

## Prologue/Background

- Sediment transport consists of suspended load and bedload (saltating particles)
- Bedload is typically difficult to measure
- Since bedload impacts cause ground motion, **can we use seismic data to infer bedload flux?**

# Quantifying Sediment Transport with Seismic Measurements



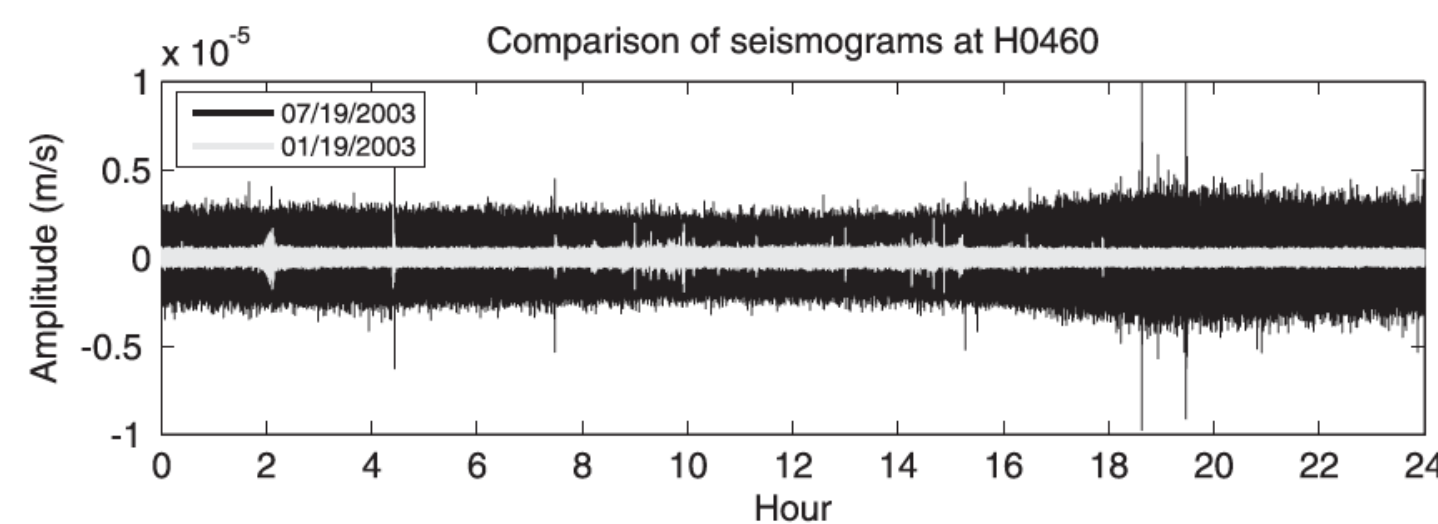
Victor C. Tsai, Brent Minchew, Michael P. Lamb, Jean-Paul Ampuero  
Seismological Laboratory, California Institute of Technology

## References

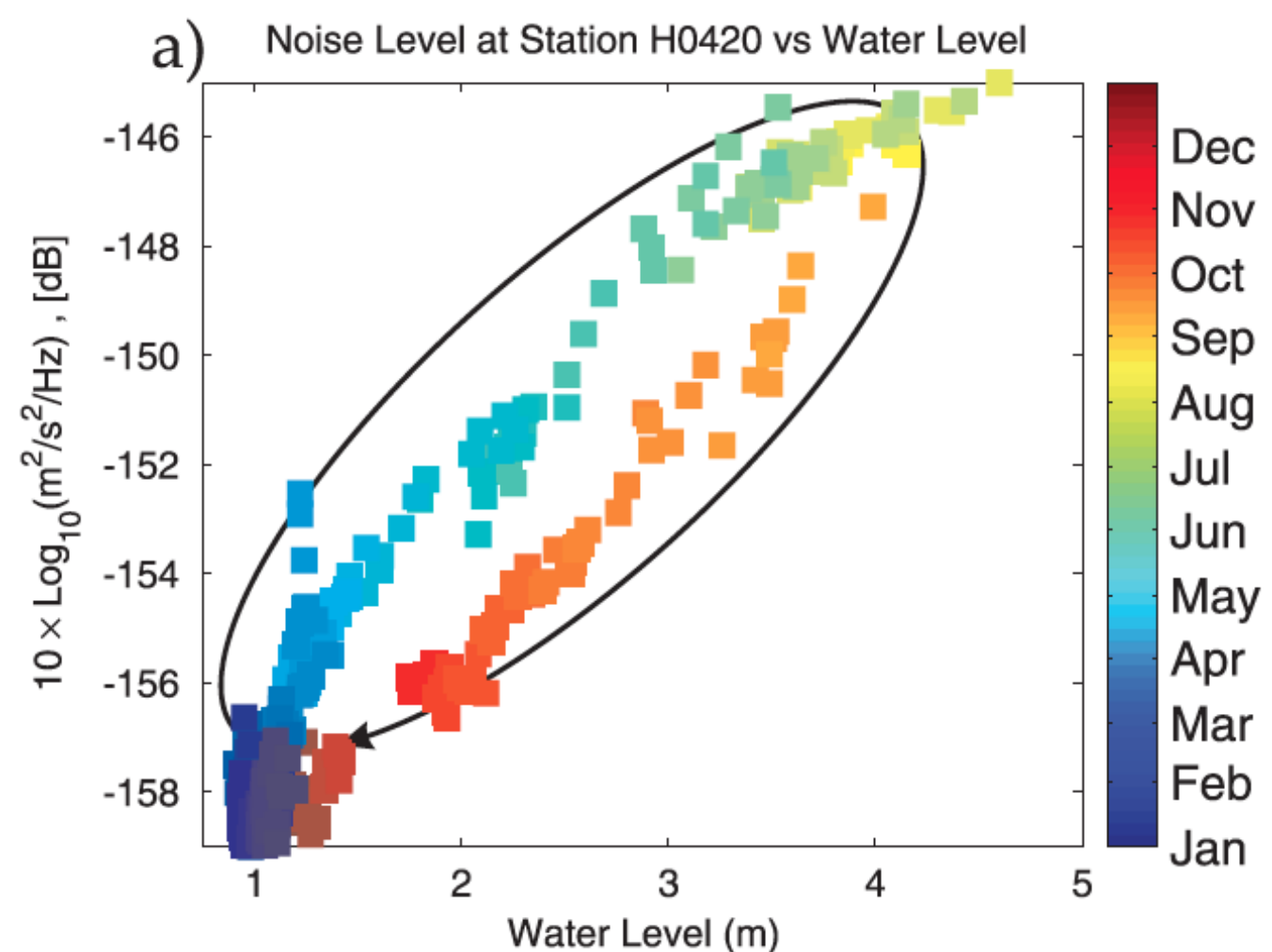
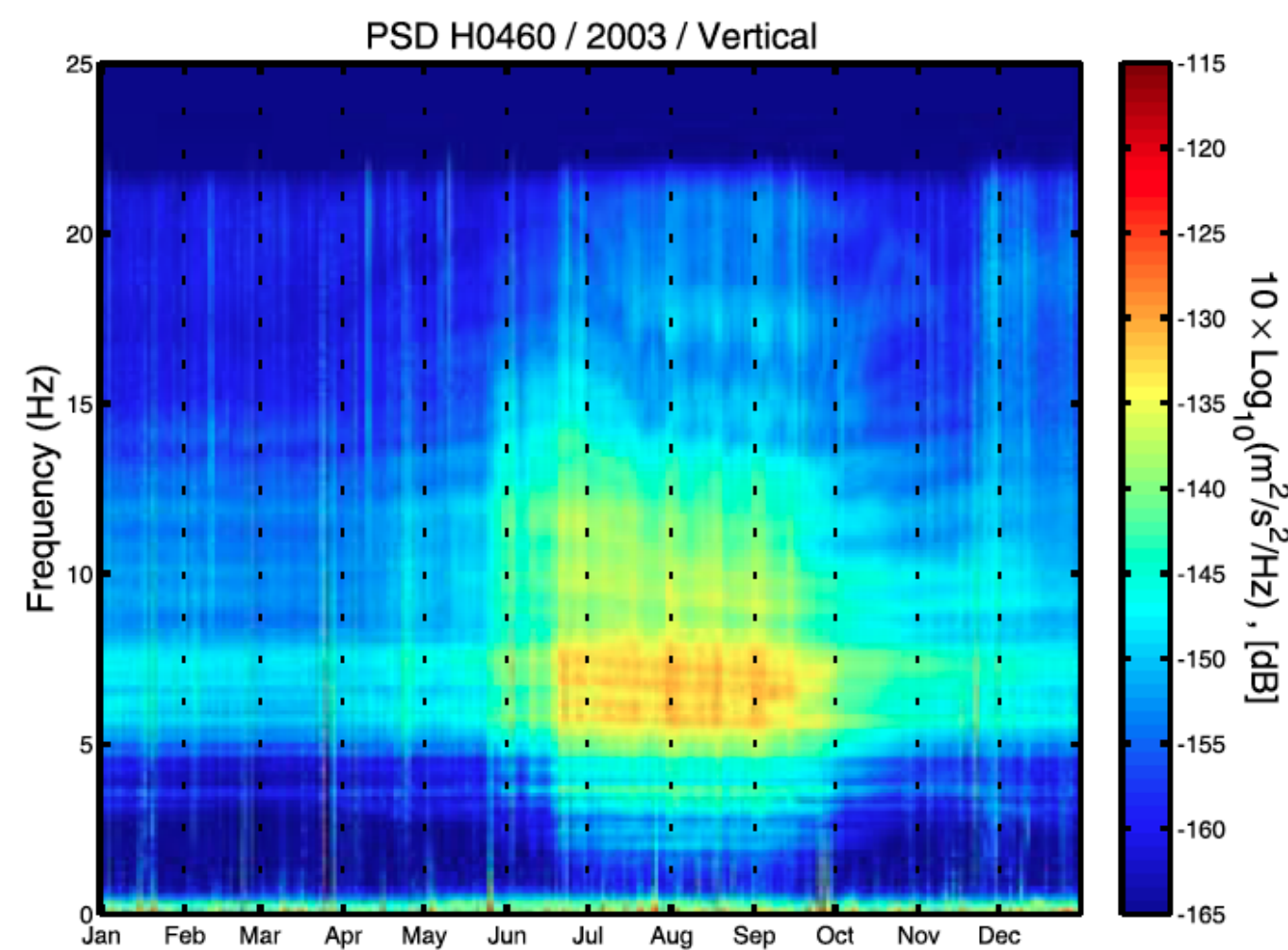
- Burtin, A., L. Bollinger, J. Vergne, R. Cattin, and J. L. Nabelek (2008), *J. Geophys. Res.*, 113, B05301.
- Tsai, V.C., B. Minchew, M.P. Lamb, J.-P. Ampuero (2012), *Geophys. Res. Lett.*, 39, L02404.

## Observations of Burtin et al.

- Data from a seismic transect near the Trisuli River in Nepal
- (The Trisuli is one of the major trans-Himalayan rivers. It has steep slopes and large, seasonal sediment flux)



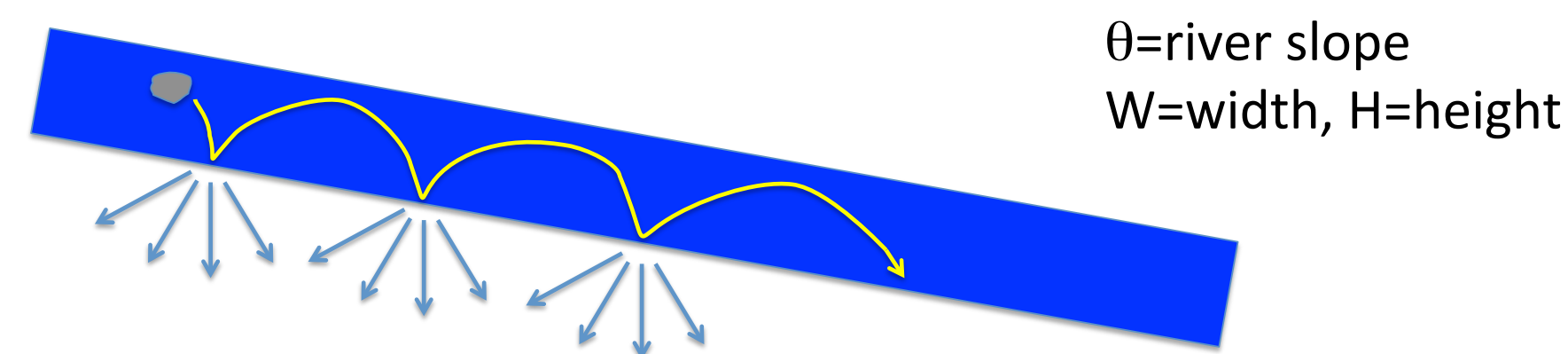
- Power spectral density (PSD) plots, showing noise energy



- Observations from Nepal suggest that **bedload contributes significant ambient noise energy**

## Modeling Seismic Noise from Sediment Transport

- To use seismic observations to solve for sediment flux, **we need a physical model to relate sediment flux to seismic noise**
- We construct a simple idealized model based on vertical particle impacts**, as schematically drawn below



- Impact rate** depends on flow speed, hop time, ...

$$\frac{dn}{dt} = \frac{CWq_b \bar{w}}{VU_b H_b} \quad (q_b = \text{sediment flux, } w = \text{settling velocity, } U_b = \text{horiz. particle speed, } H_b = \text{bedload height})$$

- Impact strength** depends on particle impulse (mass, speed)

$$F \Delta t \approx 2mw_i \quad (m = \text{particle mass, } w_i = \text{vertical impact speed})$$

- Ground motion assumed to mostly be **Rayleigh waves** for which we assume a simple analytical Green's function

$$|G(f, x; x_0)| \approx \frac{k}{8\rho v_c v_u} \sqrt{\frac{2}{\pi kr}} \cdot e^{-\pi fr/(v_u Q)}$$

- With these assumptions, **noise power  $P_v$  can be predicted:**

Integrate all impacts  $\rightarrow$

$$P_v(f; D) \approx \frac{dn}{dt} \frac{\pi^2 f^3 m^2 w_i^2}{\rho^2 v_c^3 v_u^2} \chi(\beta)$$

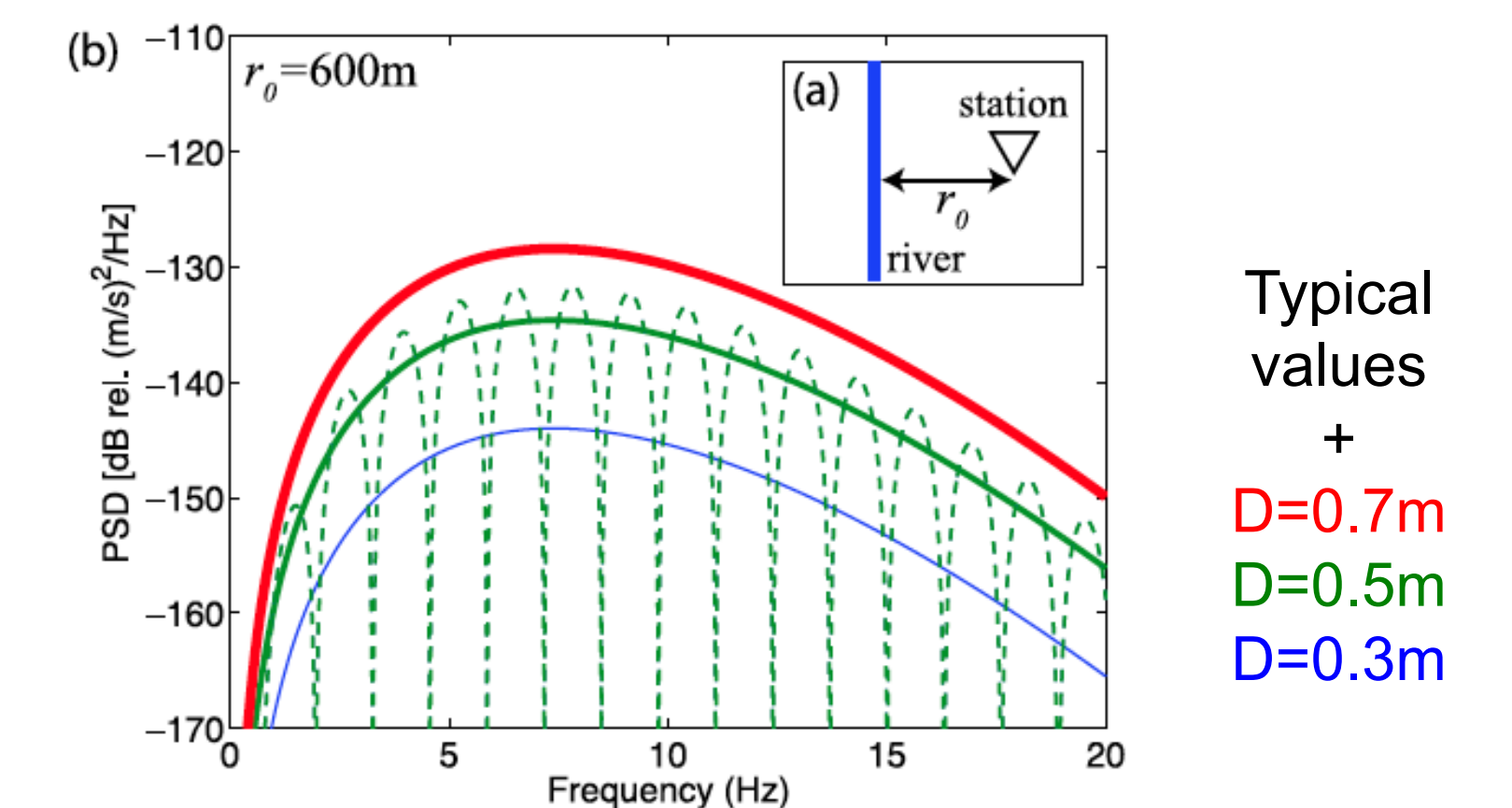
$$\chi(\beta) = \dots$$

- This expression assumes a linear river, and is frequency ( $f$ ) and grain size ( $D$ ) dependent, and also depends on a number of other parameters like phase velocity ( $v_c$ ) and attenuation ( $Q$ )

- Note dependence on  $m^2$  and  $w_i^2$ .** This implies large particles and faster impacts are more important!

## Model Results/Predictions

- Given  $D, q_b, \theta, H, W, v_c, \dots$   
**We predict seismic power spectral density (PSD)**

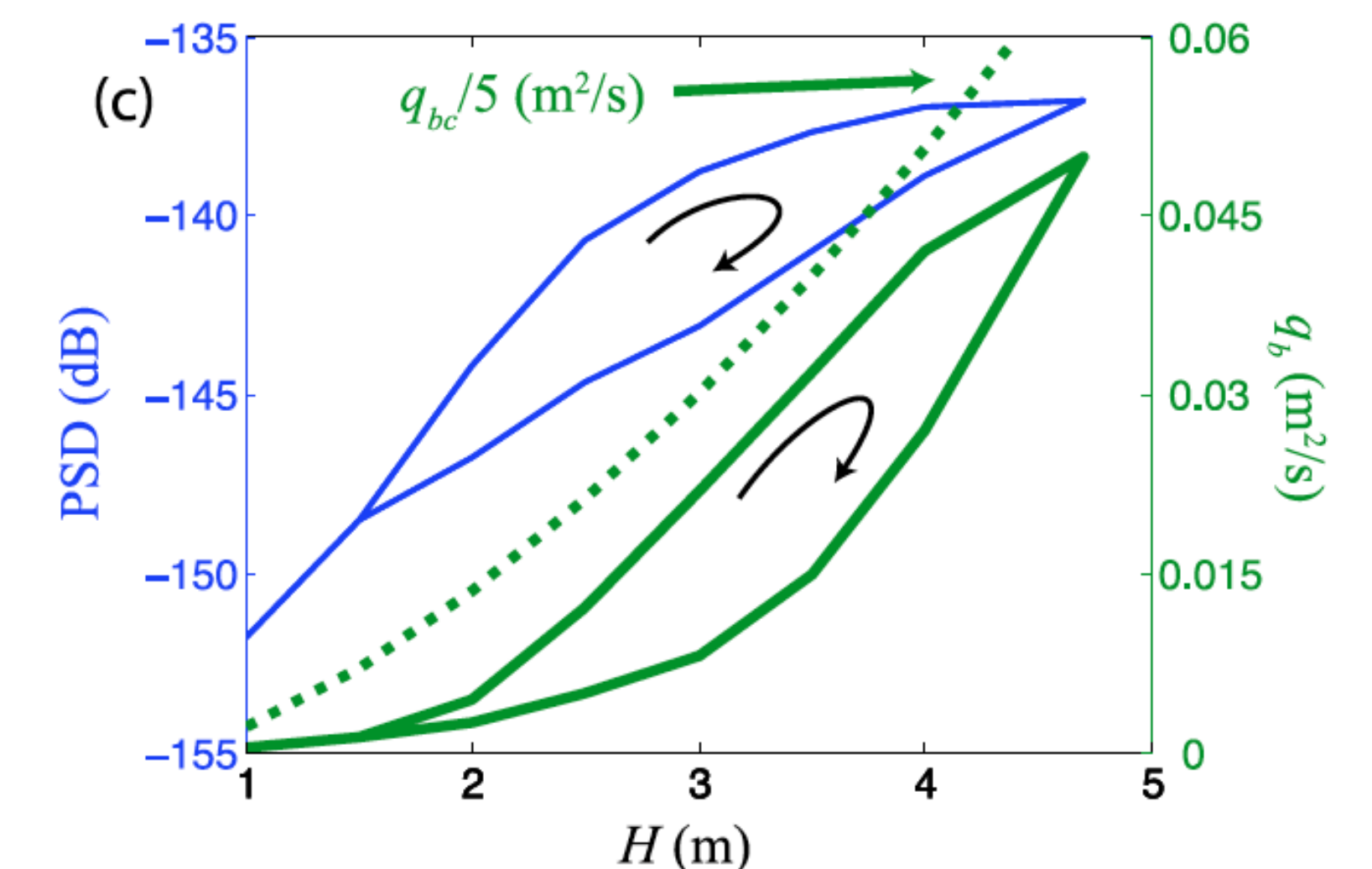


- Solid curves smoothed. Dashed curve includes effect of hop time on frequencies. Predictions are for single grain sizes noted at right

- Since we now have a forward model, we can invert seismic for  $q_b$ . In other words:  
Given  $D, PSD, \theta, H, W, v_c, \dots$

**We predict bedload sediment flux ( $q_b$ )**

- Need to assume grain size distribution, and currently we assume water flow noise is insignificant (may not be a good assumption over falling limb)



- Green solid curve is our prediction of  $q_b$  from the data of Burtin et al. (shown at left, and idealized as the blue curve)

- We predict bedload sediment flux!**
- But are we right?? (Future work... calibration, etc.)