



Abstract

Recent advances in observations have highlighted fault behavior over different temporal scales, from rapid seismic motion to slower postseismic slip, and complex interseismic behavior, including episodic phases of accelerated slip. Many observations document complex spatial variations of faulting encompassing steady and oscillatory creep, unstable sliding, and more recently deep tremors. The difficulty to simulate realistic earthquake cycles in space and time can be attributed to the large and essentially unknown spatial variations in fault properties, a limited detailed history of past events and to the nonlinear friction law that controls fault dynamics. However, laboratory experiments and theoretical developments provide an increasingly realistic and detailed understanding of the fault physics, offering the basis for an extrapolation to natural conditions. Here, we bridge the gap between observations and fault physics by developing the first model of the earthquake cycle to explain many puzzling, yet robust observations of the crustal dynamics at Parkfield.

The Parkfield segment of the San Andreas Fault (SAF) is an ideal natural laboratory to model the earthquake cycle because of its rich behavior and the presence of an unprecedented array of observations, including the SAFOD drilling site and modern seismic and geodetic networks. Despite the similarities between the previous events and their short recurrence times (from 12 to 32 years for 5 events until 1966), the latest rupture of 2004 defied the statistics by taking place a decade later than anticipated and initiated to the south to propagate northward, contrarily to all previous events.

We present the first physical model of the full earthquake cycle at Parkfield that includes all stages of the deformation and reconciles many important observations while incorporating realistic aspects of fault dynamics. Our model is compatible with the contemporary relative motion between the North-American Plate and the Pacific Plate across the SAF and gives rise to a sequence of Mw 6 earthquakes that can explain the observed variability of hypocenters and reproduce the geodetic observations of surface deformation in the co- and postseismic periods associated with the 2004 event. Our study introduces a methodology to integrate seismological and geodetic observations into a physical model of the earthquake cycle that can help better understand and mitigate seismic hazard around active faults.

Our interpretation of the microseismicity circumscribing the SZ reconciles various observations of the earthquake cycle at Parkfield such as the inferred slip distributions of the 2004 earthquake from seismic and geodetic data, the secular GPS velocity, and our best estimate of the slip distribution for the 1966 earthquake. We terminate the SZ to the south close to the nucleation site of the 2004 event assuming that the hypocenters of moderate-size earthquakes are located in areas of stress concentration which can occur between stable and unstable friction.



A physical model of the co- and postseismic deformation

Our fully dynamic model of the earthquake cycle offers an excellent fit to the coseismic GPS displacement vectors and the daily time series of postseismic GPS displacement.



Under the hood of the earthquake machine: towards predictive modeling of fault slip

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The special role of microseismicity at Parkfield: marking the locked/creeping transition

A reference model for the sequence of Mw 6 earthquakes at Parkfield

We obtain a long and rich history of fault slip with spontaneous nucleations and ruptures of earthquakes of magnitude ranging from Mw 2 to 6. The minimum size of simulated earthquakes in our models is bounded by finite computational resources and smaller earthquakes could be obtained on a finer numerical mesh. Some snapshots illustrative of the modeled fault dynamics are shown below for the case of a typical Mw6 earthquake, for which most coseismic slip occurs in the area circumscribed by microseismicity.



$$V = 2V_0 \,\theta^{-b/a} \sinh \frac{\tau}{a(\sigma_n - p)}$$